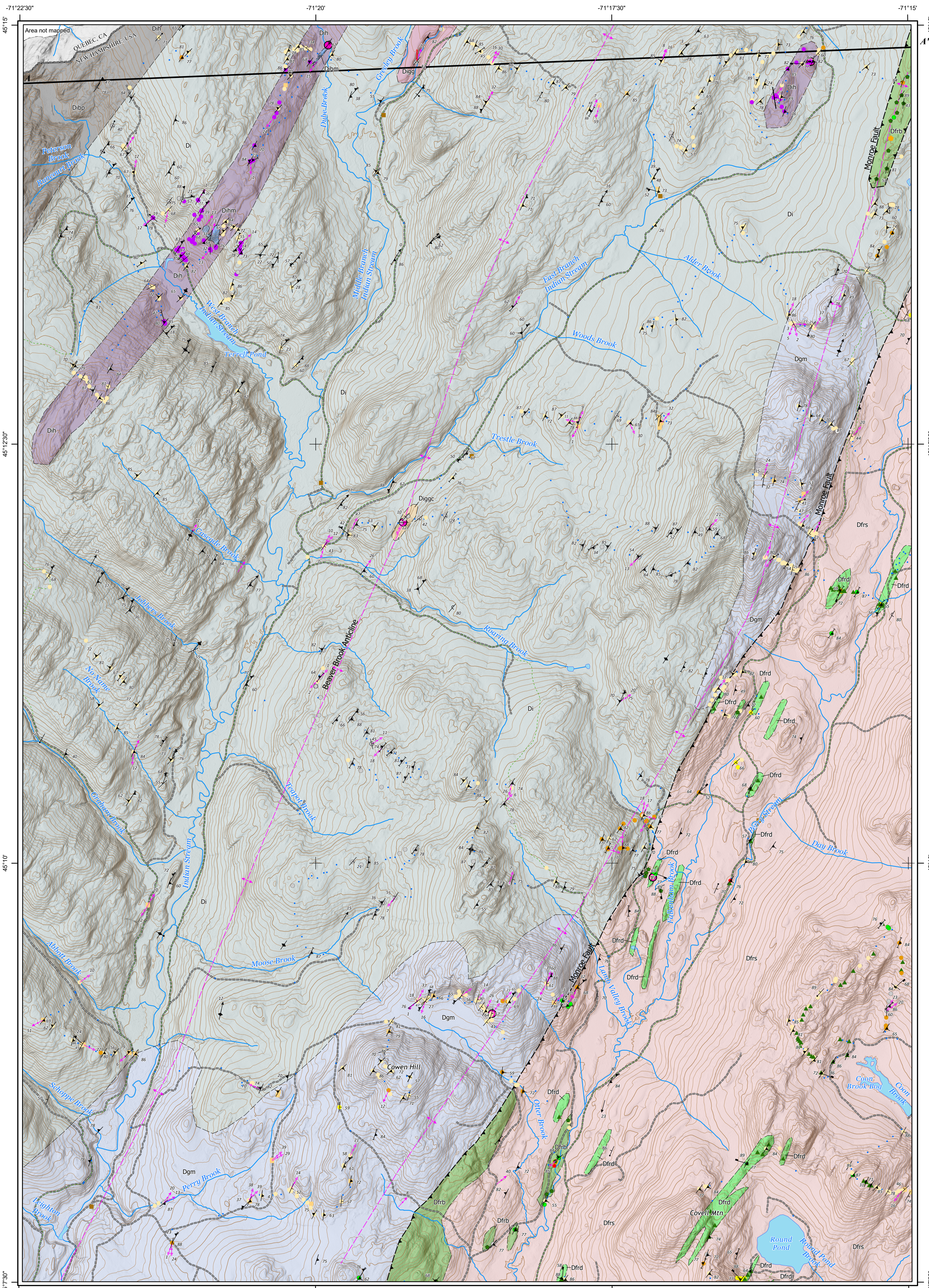


Bedrock Geologic Map of the Cowen Hill 7.5' Quadrangle, New Hampshire, 2022



DESCRIPTION OF MAP UNITS

IGNEOUS AND META-IGNEOUS UNITS Mesozoic White Mountain Plutonic-Volcanic Suite

Mzf Felsic dike — Mzf — Light red brown to golden weathering 1-1.2m thick, dark gray fine-grained micro-porphyrific to sugary apite dike cross-cutting thick-bedded Devonian graywacke in the Indian Stream Canyon (Middle Branch) — see Cross-Section A-A'. Feldspars are dominant, but much altered to form a fine-grained sericitic matrix. Potential sources are the Jurassic Monadnock Mountain (Vt) and the Cretaceous Mont Mégantic (Quebec).

Early Devonian Frontenac Formation meta-igneous units

The relationships between the Frontenac intrusive and extrusive metaigneous rocks can be complex and subtle. An outcrop with abundant pillow meta-basalts or lapilli tuffs that define extrusive activity, can also include coarse-grained metabasite. Contacts between coarse-grained and finer-grained units may not have an apparent chill zone. Two metaigneous Frontenac members were observed in the Cowen Hill Quadrangle.

Dfrd Metadiabase — Metadiabase: Dominantly coarser grained meta-diabases to metagabbros that are sometimes associated with meta-lapilli tuffs, pillow basalts, meta-volcaniclastic sediments and other extrusive features. Compositions are similar to the meta-basalt unit (Dfrb). Outcrops range from thin silts to thick bodies that form 30-40 m cliffs that extend up to 1-2 kilometers. It was not possible to map all occurrences due to access issues. Lapilli meta-tuffaceous sediment (DFR_MS) adjacent to metabasite bodies yielded ages of 408 ± 6 and 415 ± 5 Ma.

Dfrb Meta-basalt — Meta-basalt: Dominantly meta-basalts, common pillows, possibly associated exhalative horizons in thicker units just east of the Monroe Line. The pillows range from small, ≤ 30 cm to 2 m in length; tops can sometimes be determined. Compositions range from basalts and trachybasalts to basaltic andesites and basaltic trachyandesites. It seems likely that the meta-basalt sequence is not much older than nearby meta-volcaniclastic tuffaceous sediments with maximum depositional ages of 408 ± 6 and 415 ± 5 Ma.

METASEDIMENTARY ROCKS

Middle to Early Devonian

Di Ironbound Mountain Formation — black, metamudstone, slate, and phyllite, variably graphitic and pyritic, commonly spotted with either weathered pyrite and/or ankerite; interfingers with black metasilstones. Thin interlayers of fine-grained, light gray to rusty metasilstones preserve grading. Where graded beds are preserved, delicate crossbedding may be seen on cleavage surfaces cutting bedding at a shallow angle. Metasilstone channels occur primarily in the Hall Stream and Graded Graywacke members, but an outcrop of a stacked channel complex was also observed (Diggs). Correlative with the Ironbound Mountain Formation in northwestern Maine (Marvinney et al., 2001). Correlative with Meetinghouse Slate member of the Gile Mountain Formation in Vermont and in part with the Ludger and Lac Drolet members of the Compton Formation in Quebec.

Dibp Black Phyllite — Metamorphosed black, mudstone, slate and phyllite member has very little siltstone and represents deposition in a quiescent euxinic environment. Dibp transitions to siltier facies to the east and is correlative with the Ludger member of the Compton Formation in southern Quebec. No known fossils or age dates.

Dihm Halls Stream Grits — Discontinuous lenses up to a few km long and few hundred meters wide. May occur in layers a few cm to tens of meters thick and is characterized by abundant feldspar grains (laths) occurring in black phyllite or metasilstone matrices, including siltstone channels. In some locations, composed dominantly of feldspars and grades from thin-bedded unit (10-15 cm thick) to a massive unit without obvious bedding (Dihm). Traverses along the Dih unit north of Terrell Pond show highly variable interlayering or interfingering of the feldspar-rich units with black phyllites and siltstones. Coarser grit is composed of randomly oriented, euhedral, frequently fractured twinned plagioclase, variable amounts of clear quartz grains, and rare rock fragments in slaty/silty matrix. Maximum depositional ages vary between 413 to 425 ± 7 Ma and may be older than the actual age (see following Dthm description).

Dihm Halls Stream Grits - massive — a sub-member that is a massive feldspar-rich unit that forms both small and large cliffs, which in places do not exhibit obvious bedding. Possible interpretation of the lack of bedding is due to its origin as a submarine debris. Present in the ridge to the north of Terrell Pond and in the Pittsburg Quadrangle. Age dating of a sample exhibited a single well-defined "maximum depositional age" of 391 ± 8 Ma. Possible origin as a very rich tuffaceous sediment.

Digg Graded Graywacke — Bands of rhythmically graded, gray weathering, fine-grained micaceous feldspathic meta-sandstone and dark gray meta-shale member of the Ironbound Mountain Formation. Beds range are typically only 0.5 cm to several meters thick (e.g., Indian Stream Canyon). Quartz, feldspar, white mica and opaque are dominant phases. Interpreted to equate with Dgg in Vermont (Ratcliffe et al., 2011) and the Lac Drolet member of the Compton Formation (Dcd) in nearby Quebec where Perrot et al. (2018) reported a maximum depositional age of 413 ± 7 Ma from U/Pb analyses of detrital zircons.

Diggc Graded Graywacke Channel Complex — Thick-bedded meta-siltstone — meta-sandstone in an apparent stacked channel complex.

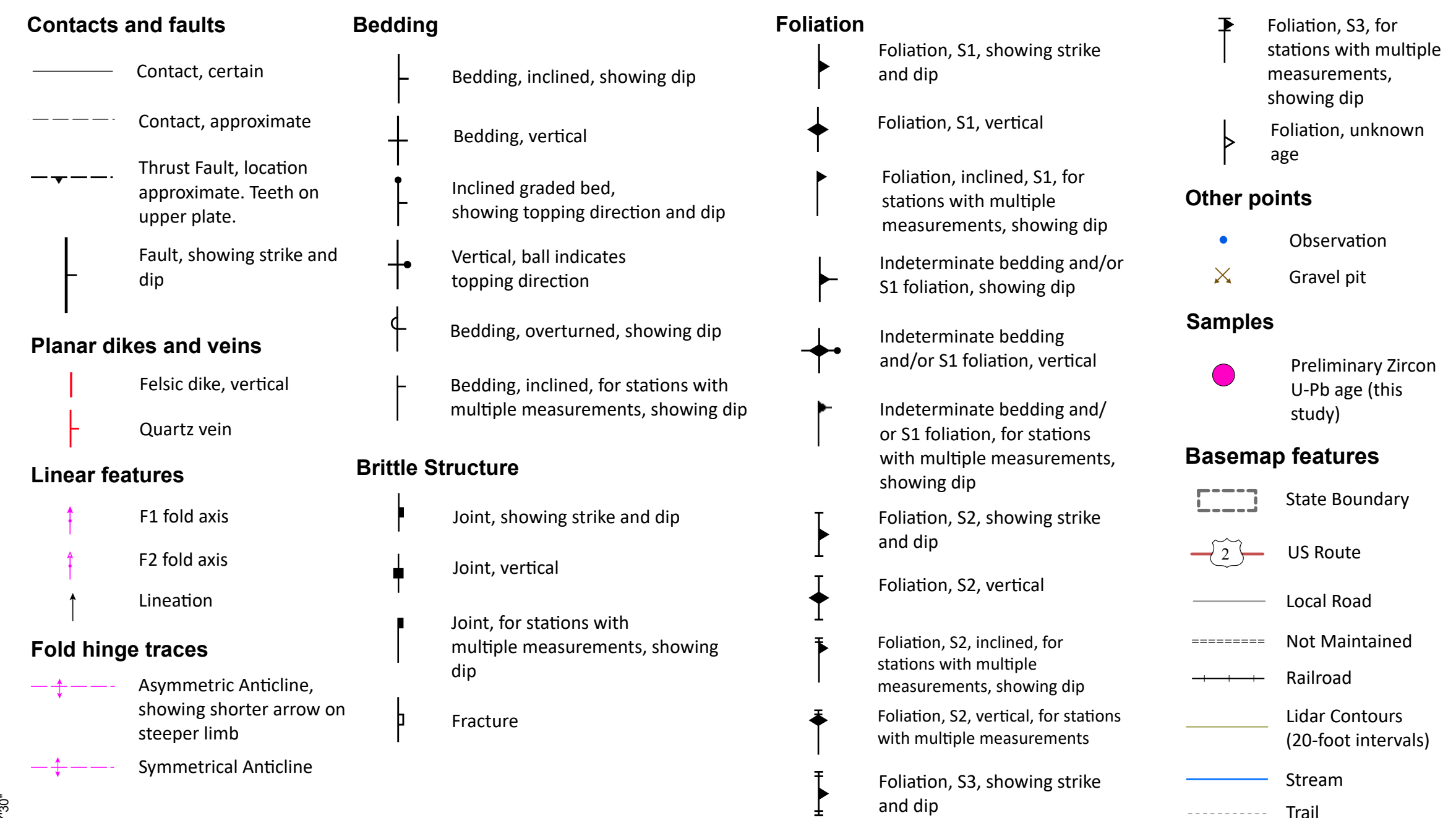
Dgm Gile Mountain Formation — Gray to light brown weathering, variably calcareous feldspathic metasediments in beds 0.5 – 3 m thick, meta-siltstone and dark gray phyllite. These units often carry ankerite and/or pyrite cubes up to 2 cm across; alteration of both minerals produces a reddish-brown weathering surface. Graded bedding is rarely preserved; brown weathering carbonate-rich lenses are present locally. Two detrital zircon samples (Lake Francis Quad) yielded maximum depositional ages of 413 ± 8 Ma and $411 - 414 \pm 10$ Ma respectively. Just north of Cowen Hill, detrital zircon analyses yielded several young age clusters: 388 ± 8 , 411 ± 6 and 426 ± 7 Ma and a crystallization age (basaltic andesite) yielded an age of 40 ± 6 Ma. Correlated the Milan member (maximum depositional age of 413 ± 7 Ma along the Quebec-NH border - Perrot, 2018) of the Compton Formation and the Gile Mountain Formation (Dggs) in nearby Vermont (maximum depositional age of 401 ± 6 Ma near St. Johnsbury (Perrot, 2018)).

Dfrs Metasediments of the Frontenac Formation Frontenac metasediments consist of metasilstones, sparse metasediments, phyllites and a wide range of volcaniclastic metasediments. Homogenous metasilstone bed thicknesses can exceed 1-2 m thick. In some places, there are graded beds that preserve Bouma couplets, which occasionally provide top directions, and rare calcareous lenses up to a meter in length. A distinct characteristic of much of the Frontenac metasediments is the presence of sufficient chlorite to provide a green color that is uncommon in the Gile Mountain metasediments.

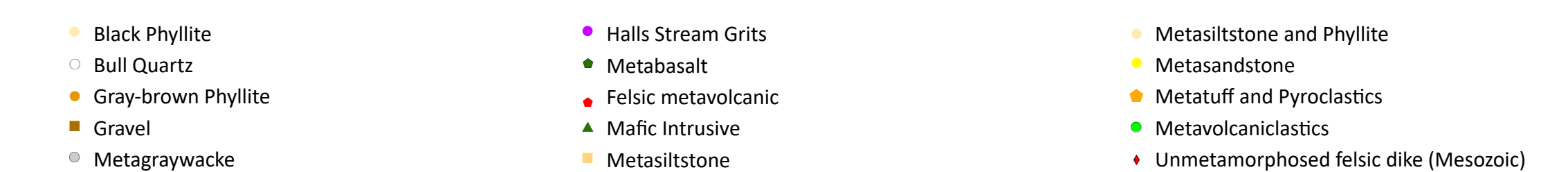
Two U-Pb zircon analyses from meta-volcaniclastic sediments / lapilli tuffs in the Lake Francis Quadrangle yielded maximum depositional ages of ca. 408 ± 6 Ma and 415 ± 6 (new) Ma. A third analysis of very-fine grained metasediment recovered a limited number of zircons yielded a maximum depositional age of 443 ± 16 Ma that is likely significantly older than the depositional age. In the Cowen Hill Quadrangle, a metasediment yielded a maximum depositional age of 434 ± 4 Ma. Comparison of the age distributions of zircon grains in the Frontenac and the Gile Mountain metasediments shows that most Frontenac metasediments have a unique peak around 580 Ma, suggesting a different sediment source at that time.

Dwrs Waits River Formation — Is shown at depth in Cross-Section A-A'. Interbedded brown weathering metasilstone, dark gray phyllite, and brown lenses of meta-limestone and thin, intensely foliated dark gray to black limy metasilstone as along West Road, Clarksville, NH. Previously interpreted as a calcareous member of the Frontenac Formation (Lyons et al. 1997). Correlated with parts of the Ayers Cliff Formation of Quebec (Perrot et al., 2018). Hatch (1963) suggested that the Waits River could be a calcareous facies of the Gile Mountain. Perrot (2019) presented detrital zircon age dates for the Waits River in northern Vermont near St. Johnsbury, indicating a maximum depositional age of 406 ± 9 Ma. McWilliams et al. (2010) determined two ages in Vermont: 418 ± 7 Ma and 415 ± 2 Ma. Alekoff and Karabinos (1990) reported an age of ca. 423 ± 4 Ma on a felsic dike that crosscut the Standing Pond Volcanic Member.

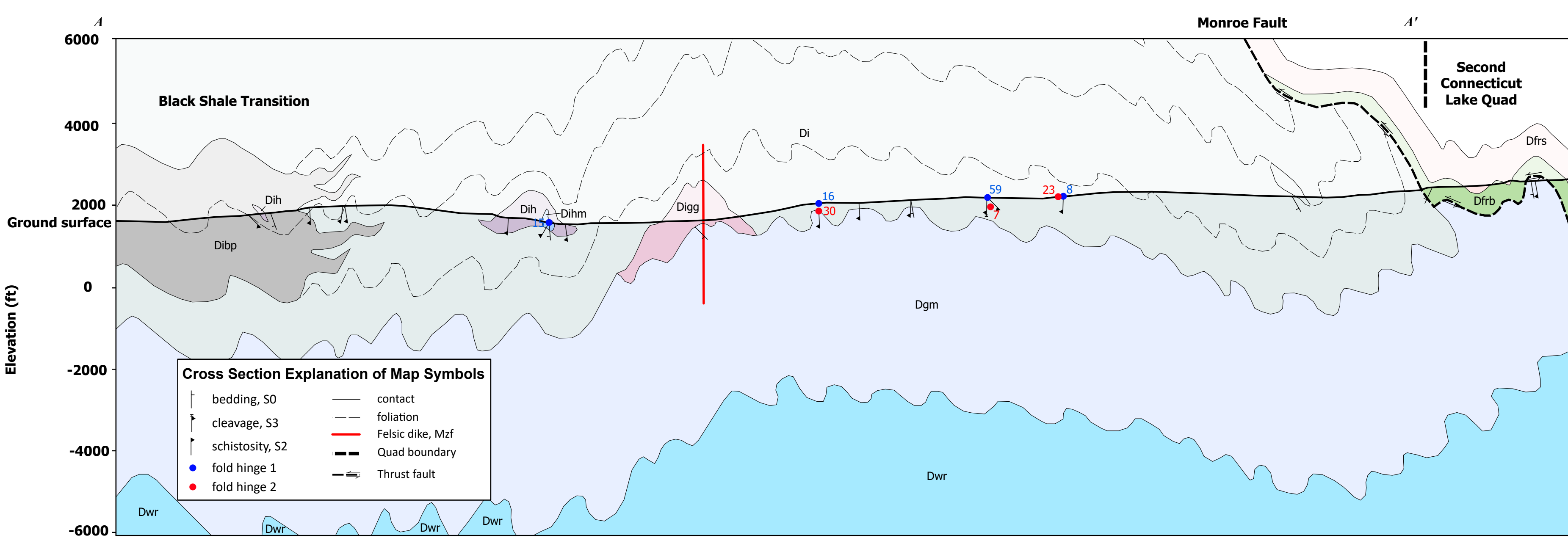
EXPLANATION OF MAP SYMBOLS



EXPLANATION OF LITHOLOGY POINT SYMBOLS



Interpretive Cross Section A - A' (No Vertical Exaggeration)



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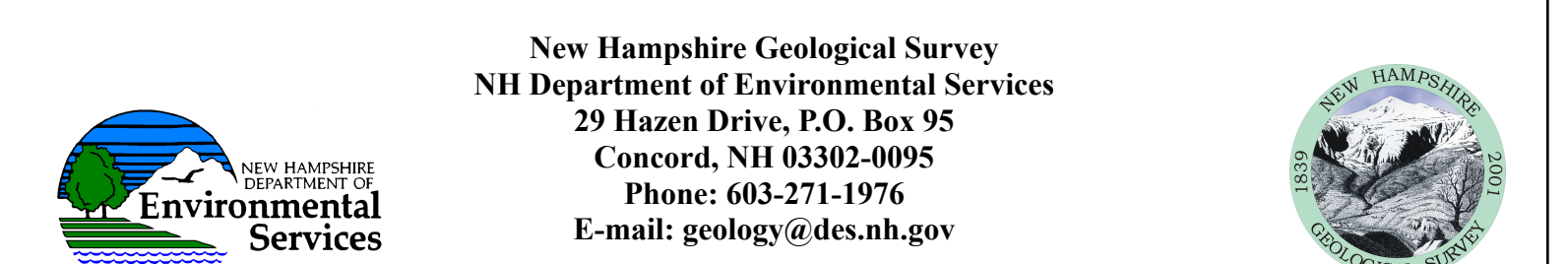
Geology by D.R. Converse, W.A. Bothner, C.E. Jahring, II, and P.S. Koch, 2022
Digital Compilation by D.R. Converse, and J.A. Keeley, 2022

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Maps and map pamphlets can be found at <https://www.des.nh.gov/land/geology>

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